Monopole source applied to volcanic eruptions

### Jeffrey Johnson Volcano Acoustics Workshop IAVCEI Kagoshima 2013

# Simple acoustic monopole source\* applied to volcanic eruptions:

### Simplifications often invoked:

- Source is simple acoustic (i.e., point source or compact)
  -> characteristic wavelength is >> vent region (ka<<1)</li>
- Propagation is simple acoustic (i.e., omnidirectional)
  -> raypaths spread radially
- Sound recorded is purely far-field term (kr>>1)
  -> pressure falls off as 1/r
- Propagation is linear
  - -> linear relation between stress and strain (no shock)

### Assumptions lead to:

 Recorded pressure is proportional to volumetric acceleration of the atmosphere

### Acoustic monopole source (sinusoid): $\delta p(r, t) = i\omega \rho_0 \frac{Q(t-r/c)}{\Omega r} e^{i\omega(t-r/c)}$

 $\delta p(r,t)$  is the excess pressure (in Pa)

r is the distance (m)

c is the sound speed (m/s)

 $\rho_0$  = density of air (kg/m<sup>3</sup>)

Q is the source strength (volumetric flux in m<sup>3</sup>/s)

 $\Omega$  is solid angle (4 $\pi$  for whole space)

# In terms of volume Flux: $\delta p(r, t) = \rho_0 \frac{\dot{Q}(t-r/c)}{\Omega r}$ where $\dot{Q} = \ddot{V}$ (volume acceleration)

 $\rho_0$  = density of air (kg/m<sup>3</sup>)  $\dot{Q}$  is 'alternative' source strength (volumetric acceleration in m<sup>3</sup>/s<sup>2</sup>)

### Acoustic monopole source (sinusoid): $\delta p(r, t) = i\omega \rho_0 \frac{Q(t-r/c)}{\Omega r} e^{i\omega(t-r/c)}$

 $\delta p(r,t)$  is the excess pressure (in Pa)

r is the distance (m)

c is the sound speed (m/s)

 $\rho_0$  = density of air (kg/m<sup>3</sup>)

Q is the source strength (volumetric flux in m<sup>3</sup>/s)

 $\Omega$  is solid angle (4 $\pi$  for whole space)

# In terms of volume Flux: $\delta p(r, t) = \rho_0 \frac{\dot{Q}(t-r/c)}{\Omega r}$ where $\dot{Q} = \ddot{V}$ (volume acceleration)

 $\rho_0$  = density of air (kg/m<sup>3</sup>)  $\dot{Q}$  is 'alternative' source strength (volumetric acceleration in m<sup>3</sup>/s<sup>2</sup>)



























More on this @: Gerst et al. (2013) The first second of volcanic eruptions from the Erebus Volcano lava lake, Antarctica - Energies, pressures, seismology, and infrasound, Journal of Geophysical Research, V. 118, 1-23

PRGME



# Acoustic Monopole Source $\delta p(r, t) = \rho_0 \frac{\dot{Q}(t-r/c)}{2\pi r}$

 $\delta p(r,t)$  is the excess pressure (in Pa)

r is the distance and c is the sound speed

 $\dot{Q}$  is the source strength (volumetric acceleration in m<sup>3</sup>/s<sup>2</sup>)

 $\Omega = 2\pi$  for hemispherical spreading















![](_page_24_Figure_0.jpeg)

![](_page_24_Figure_1.jpeg)

![](_page_25_Figure_0.jpeg)

![](_page_26_Picture_0.jpeg)

![](_page_27_Picture_0.jpeg)

![](_page_28_Picture_0.jpeg)

Event #1: occurring Jan 002 at 14:16:46

![](_page_29_Picture_0.jpeg)

![](_page_30_Picture_0.jpeg)

![](_page_31_Picture_0.jpeg)

![](_page_32_Figure_0.jpeg)

![](_page_33_Figure_0.jpeg)

### event #1: Jan002 14:16 time elapsed: 0.2 s

- ▲ 1 m uplift
- 0.5 m uplift
- 1 0.25 m downdrop

![](_page_34_Figure_4.jpeg)

#### -10 10 20 30 0 event #1: Jan002 14:16 CAL 3 Pa time elapsed: 0.3 s DOM 3 Pa CAS ▲ 1 m uplift 1 Pa • 0.5 m uplift 1 0.25 m downdrop

#### -10 10 20 30 0 event #1: Jan002 14:16 CAL 3 Pa time elapsed: 0.3 s DOM 3 Pa CAS ▲ 1 m uplift 1 Pa • 0.5 m uplift 1 0.25 m downdrop

![](_page_37_Figure_0.jpeg)

![](_page_38_Figure_0.jpeg)

![](_page_39_Figure_0.jpeg)

![](_page_40_Figure_0.jpeg)

![](_page_41_Figure_0.jpeg)

![](_page_42_Figure_0.jpeg)

![](_page_43_Figure_0.jpeg)

![](_page_44_Figure_0.jpeg)

![](_page_45_Figure_0.jpeg)